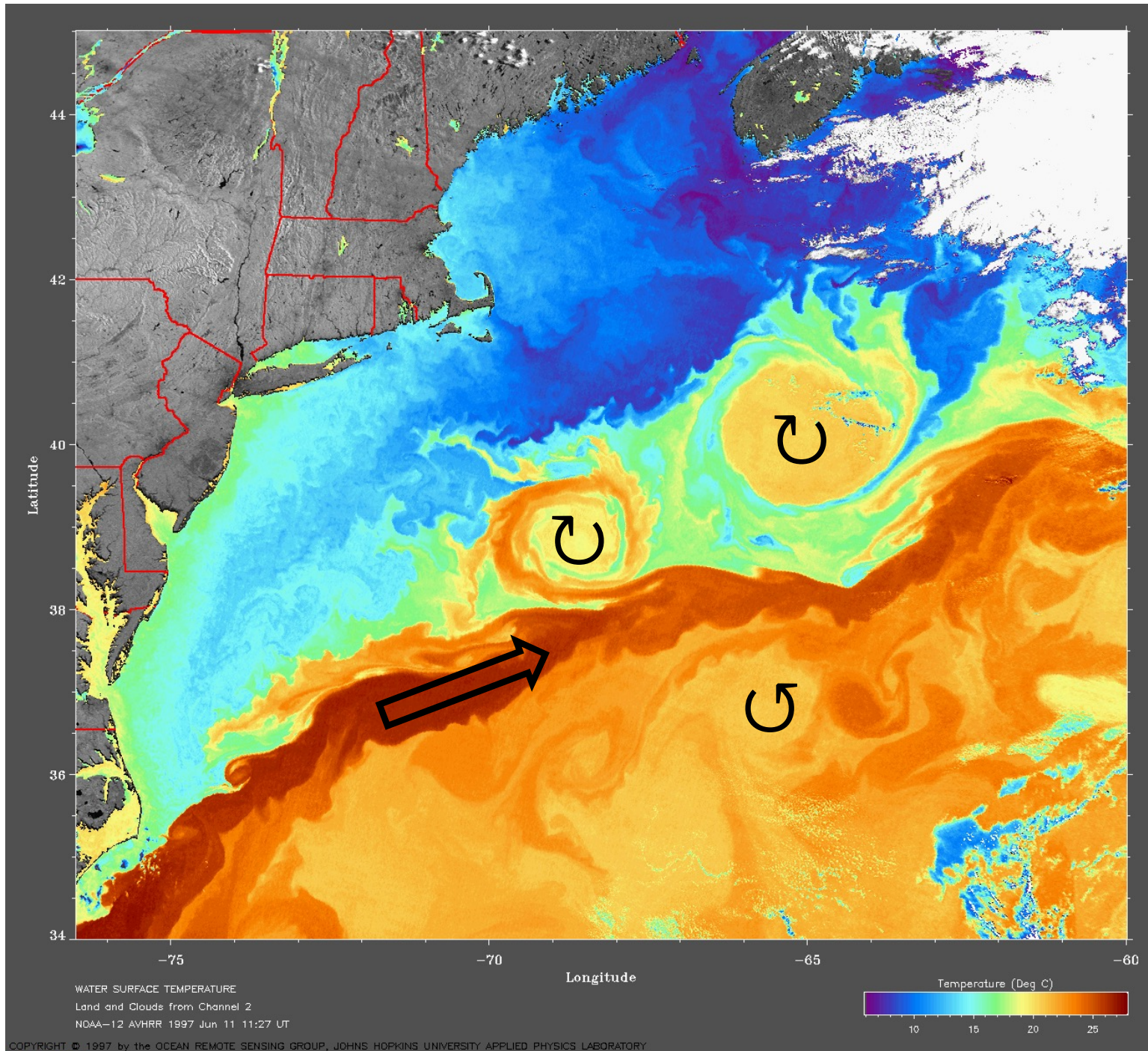


Sea Surface Temperature: Gulf Stream Eddies 1997



↺ cyclonic

↻ anti-cyclonic

↕
 $L \sim 200 \text{ km}$

$L \sim 200 \text{ km}$

$U \sim 1 \text{ m/s}$

$L/U \sim 50 \text{ hrs}$

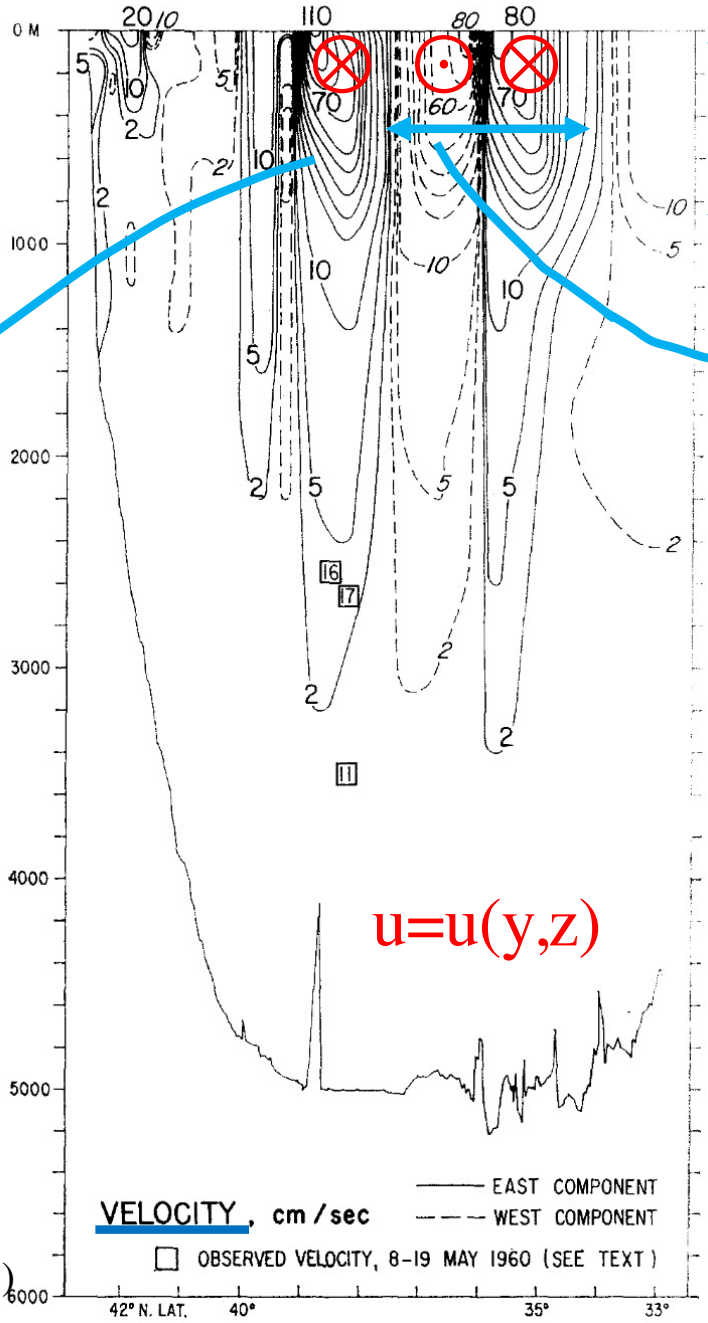
$Ro = (U/L) / \Omega$
 ~ 0.1

GFD

$L \sim 100 \text{ km}$



STATION 5907 5910 5915 5920 5926



$H \sim 1 \text{ km}$



Cyclonic eddy
or

Gulf Stream meander

Gulf
Stream

Here L is oceanic
mesoscale $\sim 100 \text{ km}$

$U \sim 0.5 \text{ m/s}$

$\partial U / \partial z \sim U / H \sim 0.5 \times 10^{-3} \text{ s}^{-1}$

0.5 hrs

$\partial U / \partial y \sim U / L \sim 0.5 \times 10^{-5} \text{ s}^{-1}$

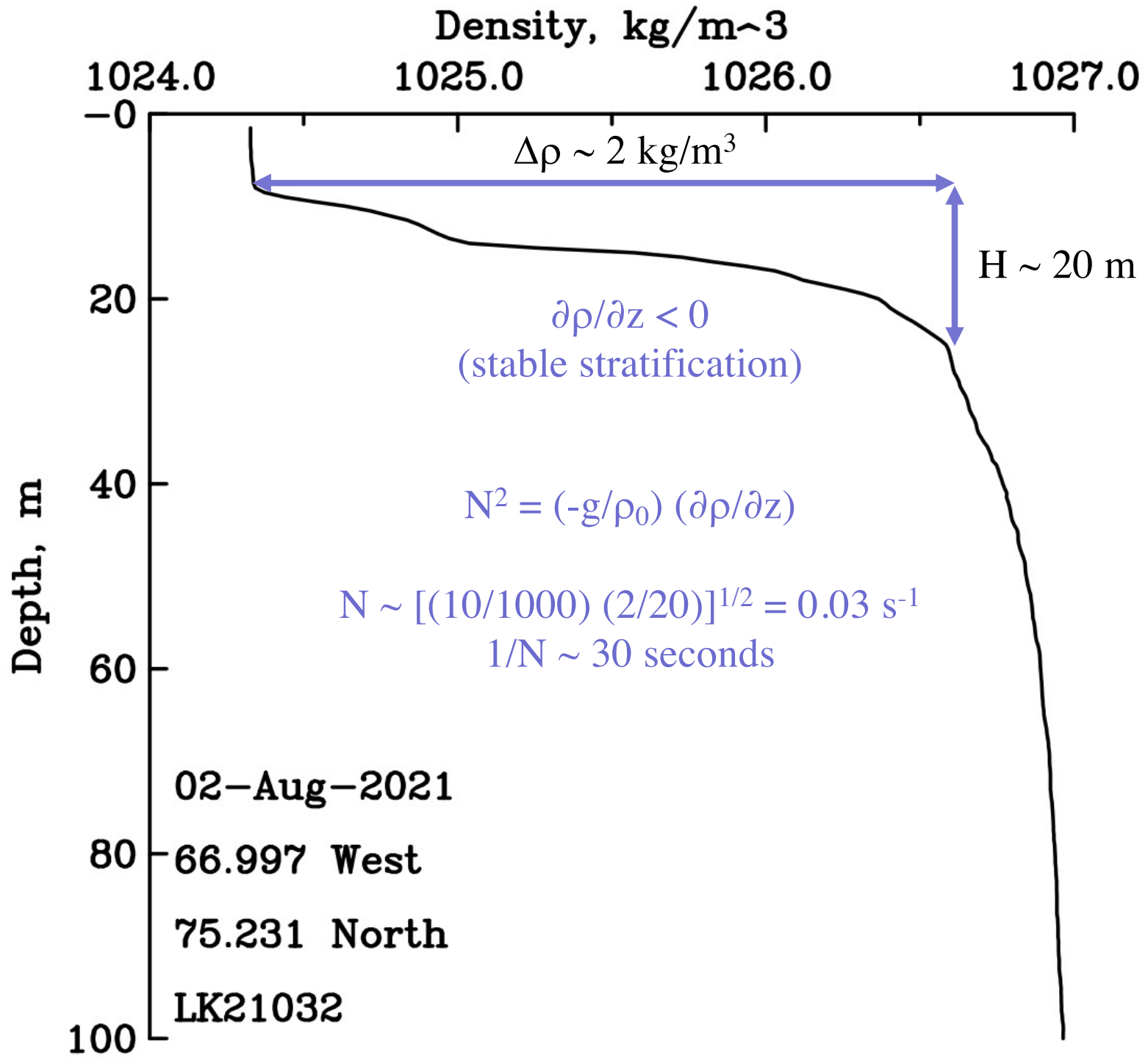
55 hrs

Fuglister (1963)

VELOCITY, cm/sec

— EAST COMPONENT
- - - WEST COMPONENT

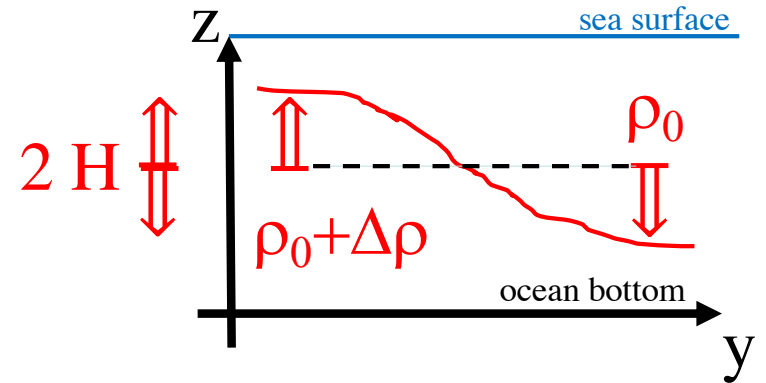
□ OBSERVED VELOCITY, 8-19 MAY 1960 (SEE TEXT)



Role of Density Stratification:

Density = Mass/Volume [kg/m³]

Froude Number = kinetic/potential energy



Kinetic Energy is $\rho_0 U^2 / 2$ per unit volume

Potential Energy

+ $(\rho_0 + \Delta\rho) g H$ of raised parcel with density $\rho_0 + \Delta\rho$ by scale height H

- $(\rho_0) g H$ of lowered parcel with density ρ_0 by scale height H

= $\Delta\rho g H$ net gain of Potential Energy (per unit volume)

Froude Number = $1/2 \rho_0 U^2 / (\Delta\rho g H)$