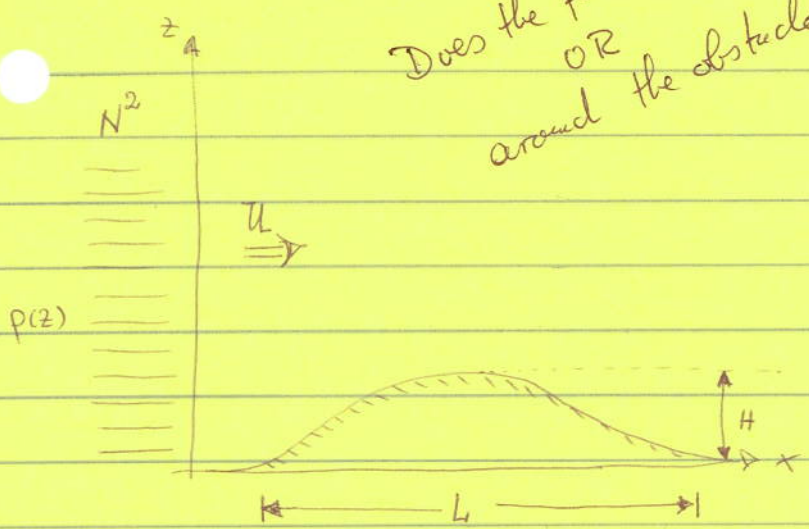


Does the flow go over OR around the obstacle?

in a density stratified fluid that has both vertical AND horizontal variations



$$T \sim L/U$$

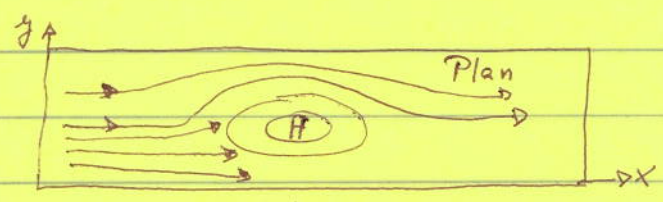
horizontal scale

$$\Delta z \sim W \cdot T \sim W \cdot L / U$$

vertical scale

but this now causes density changes

$$\Delta \rho \approx \left(\frac{\partial \rho}{\partial z} \right) \cdot (\Delta z) \sim \left(\frac{N^2 \rho_0}{g} \right) \cdot (W L / U)$$



These density variations correspond to pressure variations

Note that H is a scale height related to stratification

Scale for vertical pressure differences:

$$\Delta p \sim g H \cdot \Delta \rho$$

$$\sim g H \frac{N^2 \rho_0 (W L / U)}{g}$$

(from hydrostatic $\frac{\Delta p}{\Delta z} = \rho_0 g$ or $\Delta p \sim \Delta \rho \cdot g H$)

$\Delta z \sim H$

$$= H \frac{g \Delta \rho}{\rho_0 H} \cdot \rho_0 \frac{W L}{U} = \Delta \rho g \frac{W}{U}$$

N^2

Scale for horizontal pressure differences:

scale $\Delta p \sim \rho_0 U^2$

ageostrophic velocity

$$\frac{\partial p}{\partial x} = \rho_0 \left(u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} \right) \sim \rho_0 \frac{U^2}{L}$$

this gives

$$\rho_0 U^2 \sim \rho_0 N^2 H (W L / U)$$

$$\text{or } \frac{U^2}{N^2 H^2} \sim \frac{W/H}{U/L} = \frac{W/H}{U/L}$$

$$N^2 = -g \frac{\partial \rho}{\rho_0 \partial z} \quad \left(\frac{1}{s^2} \right)$$

$$\text{or } \left(\frac{U}{NH} \right)^2 \sim \frac{W/H}{U/L}$$

note that $(NH)^2 = -g \frac{\partial \rho}{\rho_0 \partial z} \cdot H^2$

$$\approx -g \frac{\Delta \rho}{\rho_0} H = c_i$$

vertical

$$\Delta p \sim g \cdot \Delta \rho \cdot \Delta z$$

$$\sim g \frac{\Delta \rho}{\rho_0} \frac{W}{U/L}$$

(from hydrostatic pressure perturbation $\frac{\partial p}{\partial z} = -\rho'g$)

$$\Delta z \sim W \cdot T \sim W \cdot L/U$$

horizontal

$$\Delta p \sim \rho_0 U^2$$

geostrophic velocity $\frac{\partial p}{\partial x} = \rho_0 \left(u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} \right)$

$$\sim \rho_0 U^2/L$$

$$\Delta p_{\text{vertical}} = \Delta p_{\text{horizontal}}$$

$$g \cdot \Delta \rho \frac{W}{U/L} = \rho_0 U^2$$

or

$$U^2 = \frac{g \frac{\Delta \rho}{\rho_0} \frac{W}{U/L} \cdot \Delta z}{\Delta z}$$

$$N^2 \sim \frac{-g \frac{\Delta \rho}{\rho_0}}{\Delta z}$$

$$\Delta z \sim H \sim W \cdot L/U$$

$$\sim \frac{N^2 \cdot W \cdot H}{U/L} \cdot \frac{H}{H}$$

$$\frac{U^2}{(NH)^2} \sim \frac{W/H}{U/L} \sim 1$$

$$\text{as } \frac{W}{U/L} \sim H$$

$$F_r^2 \quad \text{stiffness}$$

internal wave phase speed:

$$\text{Note that } (NH)^2 = \frac{-g}{\rho_0} \frac{\partial \rho}{\partial z} \cdot H^2 \sim \frac{g \Delta \rho}{\rho_0} H = C_{\text{internal}}^2$$

Internal Froude Number $Fr^2 = \left(\frac{U}{NH} \right)^2 = \frac{WH}{U/L}$

kinetic energy

So if $Fr \leq 1$

potential energy
due to buoyancy

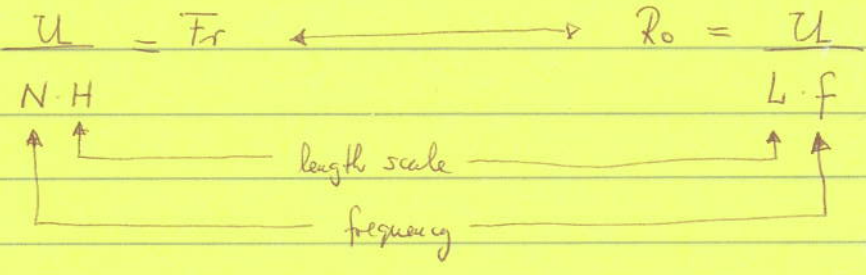
$Fr \leq 1$

stratification forces fluid $\Delta z \leq H$ vertically
or $WH \leq U/L$

and a portion of
the fluid will go around (rather than over) the obstacle
→ recall Taylor-Proudman theorem

strong stratification → large N → small Froude Number → stiff fluid

Also compare



For QG dynamics we can expand
(almost geostrophic)

QG "Quasi-Geostrophic"

$$u = u_0 + Ro u_1 + \dots$$

$$v = v_0 + Ro v_1 + \dots$$

↑
geostrophic velocity

Horizontal divergence

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = \underbrace{\frac{\partial u_0}{\partial x} + \frac{\partial v_0}{\partial y}}_{\text{geostrophic} = 0} + \underbrace{Ro \left(\frac{\partial u_1}{\partial x} + \frac{\partial v_1}{\partial y} \right)}_{\sigma(Ro) \quad \sigma(U/L)} + \dots$$

$$= \sigma \left(Ro \frac{U}{L} \right)$$

Vertical divergence

balanced by horizontal divergence

$$\frac{\partial w}{\partial z} \sim \frac{W}{H} = \sigma \left(Ro \frac{U}{L} \right)$$

STIFF means $\frac{W}{H} \ll \frac{U}{L}$

stiff fluid for $Ro \ll 1$
(Taylor Prandtl)

So $\frac{W/H}{U/L} = \sigma(Ro)$

analogous

But $\frac{W/H}{U/L} \sim \left(\frac{U}{NH} \right)^2 \sim \overline{Tr}^2$

stiff fluid for $\overline{Tr} \ll 1$

Thus in a strongly stratified flow (N large) ^{and} $\overline{Tr}^2 \ll 1$,
 we can expect feature similar to Taylor columns [^]
 [particles move along f/h contours]
 and horizontally blocked flows that we encountered
 before for $Ro \ll 1$

How to scale pressure?

geostrophic scaling,
page-26a, Lecture-5

$$P_v \sim \Delta p g H$$

from vertical momentum $\frac{\partial p}{\partial z} = -\rho g$

geostrophic
inertial

$$P_{h, g}^{(g)} \sim \rho_0 L U f$$

from horizontal momentum $f u = -\frac{1}{\rho_0} \frac{\partial p}{\partial x}$

$$P_{h, i}^{(i)} \sim \rho_0 U^2$$

from $\frac{\partial p}{\partial x} = \rho_0 \left(u \frac{\partial x}{\partial x} + v \frac{\partial x}{\partial y} \right)$
thus $U \sim L f$

$$1. \frac{P_{vertical}^{(g)}}{P_{horizontal}^{(g)}} = \frac{\Delta p g H}{\rho_0 L U f} = \frac{\Delta p g H}{\rho_0} \cdot \frac{U}{L f} \cdot \frac{1}{U^2}$$

$$= \frac{g}{\rho_0 H} \Delta p g H^2 \cdot \frac{U}{L f} \cdot \frac{1}{U^2}$$

$$= \frac{N^2 H^2}{U^2} \cdot \frac{U}{L f} = \boxed{\frac{R_o}{Fr^2}}$$

$$N^2 \sim \frac{\Delta \rho g}{\rho_0 H}$$

$$(NH)^2 \sim \frac{\Delta \rho g H}{\rho_0}$$

$$2. \frac{P_{vertical}^{(i)}}{P_{horizontal}^{(i)}} = \frac{\Delta p g H}{\rho_0 U^2} = \boxed{\frac{(NH)^2}{U^2}} = \boxed{\frac{1}{Fr^2}}$$

The scale for $P_{horizontal}^{(g)} \sim P_{horizontal}^{(i)}$ only for $R_o \sim 1$

$$3. \frac{P_{horizontal}^{(g)}}{P_{horizontal}^{(i)}} = \frac{\rho_0 L U f}{\rho_0 U^2} = \frac{f L}{U} = \boxed{\frac{1}{R_o}}$$

often though $R_o \ll 1$
or hydraulic pressure (inertial)
 <
geostrophic pressure (linear)

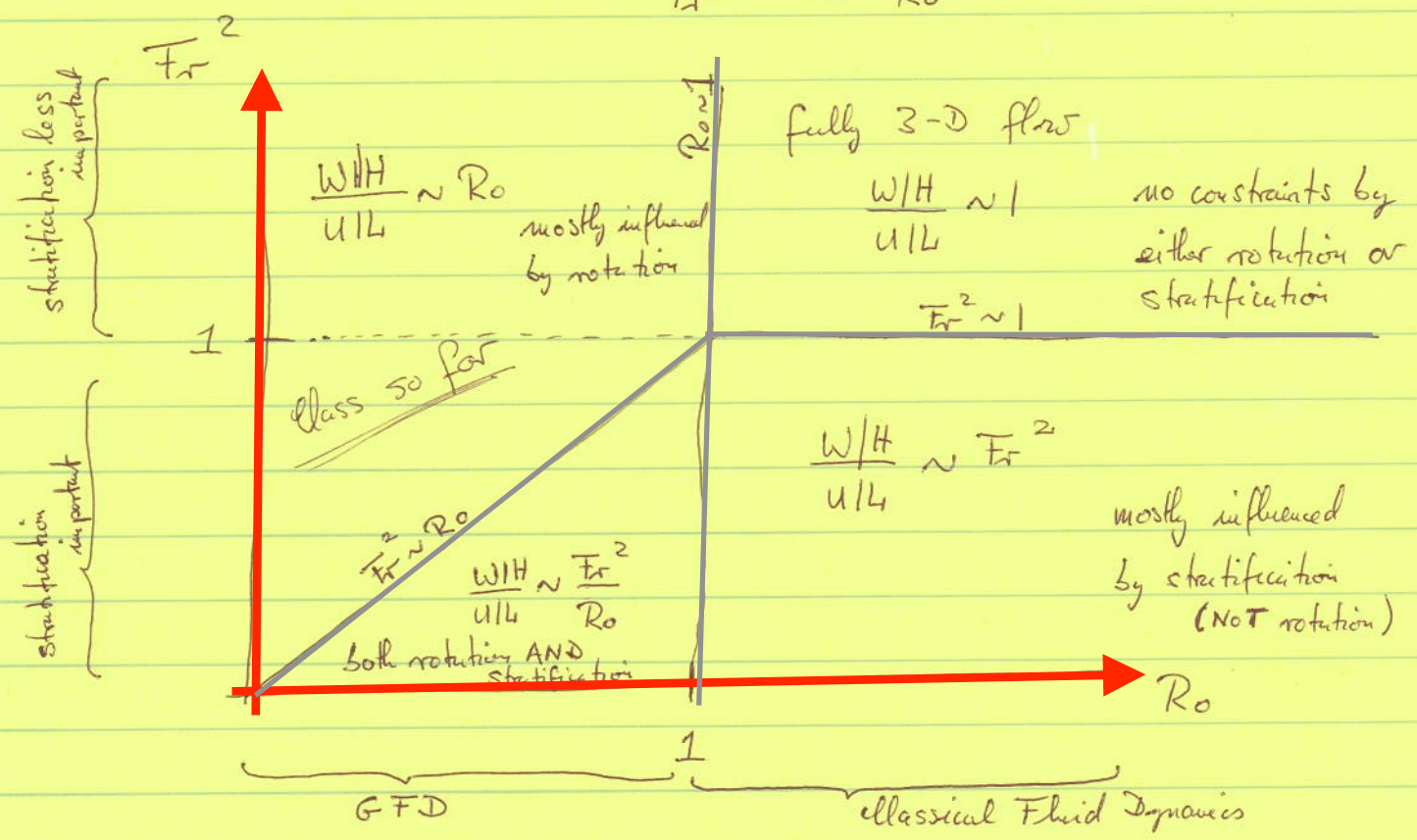
$$P_v = P_{Hg} : \Delta \rho g H = \rho_0 L \cdot U \cdot f$$

$$\text{or } \frac{\Delta \rho}{\rho_0} \frac{g}{H} \cdot H^2 = L \cdot U \cdot f$$

$$\frac{N^2 H^2}{U^2} = \frac{L \cdot f}{U}$$

$$\frac{U^2}{(NH)^2} = \frac{U}{Lf}$$

$$Fr^2 = Ro$$



$$\underbrace{\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}}_{\sim U/L} + \underbrace{\frac{\partial w}{\partial z}}_{\sim W/H} = \sigma$$